Numerical Modelling

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the anatomy of an ocean model



• Lesson 1 : [D109]

- Introduction
- Equations of motions
- Activity 1 [run an ocean model]

• Lesson 2 : [B014]

- Horizontal Discretization
- Activity 2 [Dynamics of an ocean gyre]

• Lesson 3 : [D109]

- Presentation of the model CROCO
- Dynamics of the ocean gyre
- Activity 2 [Dynamics of an ocean gyre]

• Lesson 4 : [D109]

- Numerical schemes + subgrid param
- Activity 3 [Impacts of numerics]

• Lesson 5 : [D109]

- Vertical coordinates
- Activity 3 [Impact of topography]

• Lesson 6 : [D109]

- Boundary Forcings
- Activity 4 [Design a realistic simulation]
- Lesson 7 : [D109]
 - Diagnostics and validation
 - Activity 5 [Analyze a realistic simulation]

• Lesson 8 : [D109]

Work on your projet

Presentations and material will be available at : jgula.fr/ModNum/

https://github.com/quentinjamet/ Tuto/tree/main/ModNum

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Evaluation

- The evaluation is based on a project, which consists in setting up a realistic configuration of the region of your choice, run the experiment and perform some analysis.
- Written Report due for: Jan. 29

Useful references

Extensive courses:

• MIT:

https://ocw.mit.edu/courses/earth-atmospheric-and-planetary-sciences/12-950-atmospheric-and-oceani c-modeling-spring-2004/lecture-notes/

• Princeton: https://stephengriffies.github.io/assets/pdfs/GFM_lectures.pdf

Overview on ocean modelling and current challenges:

- Griffies et al., 2000, Developments in ocean climate modelling, Ocean Modelling. <u>http://jgula.fr/ModNum/Griffiesetal00.pdf</u>
- Griffies, 2006, "Some Ocean Model Fundamentals", In "Ocean Weather Forecasting: An Integrated View of Oceanography", 2006, Springer Netherlands. <u>http://jgula.fr/ModNum/Griffies_Chapter.pdf</u>
- Fox-Kemper et al, 19, "Challenges and Prospects in Ocean Circulation Models" <u>http://jgula.fr/ModNum/FoxKemperetal19.pdf</u>

ROMS/CROCO:

- <u>https://www.croco-ocean.org/documentation/</u>
- <u>https://croco-ocean.gitlabpages.inria.fr/croco_doc/index.html</u>

 Shchepetkin, A., and J. McWilliams, 2005: The Regional Oceanic Modeling System (ROMS): A splitexplicit, free-surface, topography-following- coordinate ocean model. Ocean Modell. <u>http://jgula.fr/ModNum/ShchepetkinMcWilliams05.pdf</u>

Introduction

Master's degree 2nd year Marine Physics

Jonathan GULA gula@univ-brest.fr Quentin JAMET quentin.jamet@shom.f r • For what ocean model are used for?

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• For what ocean model are used for?

- Weather forecasts and Climate projections
- Explore different states of the ocean and earth system
- Better understand the complex dynamics of geophysical flow, from small to large scales
- Testing (idealized) mathematical models

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. . .









Can be understood using theory and simplified models

THE WESTWARD INTENSIFICATION OF WIND-DRIVEN OCEAN CURRENTS

Henry Stommel

(Contribution No. 408, Woods Hole Oceanographic Institution)

<u>Abstract</u>-A study is made of the wind-driven circulation in a homogeneous rectangular ocean under the influence of surface wind stress, linearised bottom friction, horizontal pressure gradients caused by a variable surface height, and Coriolis force.

An intense crowding of streamlines toward the western border of the ocean is discovered to be caused by variation of the Coriolis parameter with latitude. It is suggested that this process is the main reason for the formation of the intense currents (Gulf stream and others) observed in the actual oceans. JOURNAL OF METEOROLOGY

ON THE WIND-DRIVEN OCEAN CIRCULATION

By Walter H. Munk

Institute of Geophysics and Scripps Institution of Oceanography, University of California¹ (Manuscript received 24 September 1949)



of latitude



F1G. 8. Schematic presentation of circulation in a rectangular ocean resulting from zonal winds (filled arrowheads), meridional winds (open arrowheads), or both (half-filled arrowheads). The width of the arrows is an indication of the strength of the currents. The nomenclature applies to either hemisphere, but in the Southern Hemisphere the subpolar gyre is replaced largely by the Antarctic Circumpolar Current (west wind drift) flowing around the world. Geographic names of the currents in various oceans are summarized in table 3.

Munk (1950)

But reality is much more turbulent:



'Snapshot' of the global sea surface height (SSH)

Mesoscale + submesoscale

ocean colour (VIIRS – 23/09/15)

100 km

N

Mesoscale + submesoscale

100 km

ocean colour (VIIRS – 03/09/16)

A SA STORES

And HF motions like tides:



A zoo of processes and scales



Temporal and spatial scales in the ocean



Temporal and spatial scales in the ocean

"We can therefore set up seven equations independent from each other with the seven normally occurring variables. As far as it is possible to have an overview of the problem now, we must conclude that our knowledge of the laws of atmospheric processes is sufficient to serve as a basis for a rational weather prediction. However, it must be admitted that we may have overlooked important factors due to our incomplete knowledge. The interference of unknown cosmic effects is possible. Furthermore, the major atmospheric phenomena are accompanied by a long list of side effects, such as those of an electrical and optical nature. The question is to what extent such side effects could have considerable effects on the development of atmospheric processes. Such effects evidently do exist. The rainbow, for instance, will result in a modified distribution of incoming radiation and it is well known that electrical charges influence condensation processes. However, evidence is still lacking on whether processes of this kind have an impact on major atmospheric processes. At any rate, the scientific method is to begin with the simplest formulation of the problem, which is the problem posed above with seven variables and seven equations." [Bjerknes 1904]

- 7 Variables:
 - U, V, W, P, ρ , T, S (q for the atmosphere)

- 7 equations:
 - Momentum (u,v,w): $\frac{\partial \vec{u}}{\partial t} + \vec{u} \cdot \vec{\nabla} \vec{u} + 2\vec{\Omega} \times \vec{u} + g\vec{k} = -\frac{\vec{\nabla}P}{\rho} + \vec{\mathcal{F}}$
 - Equation of State (EOS):
 - Conservation of mass: (continuity)
 - Conservation of heat (of internal energy):
 - Conservation of salt: (humidity)

 $\rho = \rho(T, S, p)$ $\frac{\partial \rho}{\partial t} + \vec{\nabla} \cdot \rho \vec{u} = 0$ $\frac{DT}{Dt} = S_T$ $\frac{DS}{Dt} = S_S$

- 7 Variables:
 - U, V, W, P, ρ, T, S (q for the atmosphere)

- 7 equation
 - Momen^{*} [Navier-

But we don't know the solutions...

 $\frac{\partial \rho}{\partial t} + \vec{\nabla} \cdot \rho \vec{u} = 0$

 $\frac{DT}{Dt} = \mathcal{S}_T$

 $\frac{DS}{S} = S_S$

 $\vec{\mathcal{T}}$

- Equatio
- Conservation of mass: (continuity)
- Conservation of heat (of internal energy):
- Conservation of salt: (humidity)

Millennium Prize problems

Navier–Stokes existence and smoothness

The <u>Clay Mathematics Institute</u> in May 2000 made this problem one of its seven <u>Millennium Prize problems</u> in mathematics. It offered a <u>US \$1,000,000</u> prize to the first person providing a solution for a specific statement of the problem:^[1]

Prove or give a counter-example of the following statement:

In three space dimensions and time, given an initial velocity field, there exists a vector velocity and a scalar pressure field, which are both smooth and globally defined, that solve

the Navier–Stokes equations.

Navier-Stokes [Momentum equations]

$$\frac{\partial \vec{u}}{\partial t} + \vec{u} \cdot \vec{\nabla} \vec{u} + 2\vec{\Omega} \times \vec{u} + g\vec{k} = -\frac{\vec{\nabla}P}{\rho} + \vec{\mathcal{F}}$$

Non-linear terms = Turbulence



turbolenza by da Vinci [1507]

Big whorls have little whorls, which feed on their velocity; And little whorls have lesser whorls, And so on to viscosity. L.F. Richardson [1922]



Turbulent water jet (Re = 2300) [Van Dyke, 82]

$$\frac{\partial \vec{u}}{\partial t} + \vec{u} \cdot \vec{\nabla} \vec{u} + 2\vec{\Omega} \times \vec{u} + g\vec{k} = -\frac{\vec{\nabla}P}{\rho} + \vec{\mathcal{F}}$$

Navier-Stokes equations



Grid Discretization



Supercomputer (Curie – CEA)

One way to solve them (approximately): → Numerical modelling

Ocean modeling principle

An ocean model is simplified representation of physical processes that take place in the ocean.

The ocean is divided into boxes (finite difference, finite volume) : **discretization**

The NS equations can be solved on the grid using **numerical methods**



Ocean modeling principle

If we know:

- The ocean state at time t :
 - u, v, w, T, S, ...
- Boundary conditions : surface, bottom, lateral sides



→ We can compute **an estimated state of the ocean at time** *t+dt* using **numerical approximations of fluid dynamics equations**

Ocean modeling principle

One major Problem:

Setting the model's grid scale to the Kolmogorov length

(to resolve diffusive processes)

 $\Delta = 10^{-3}$ m over a global ocean domain of volume 1.3×10^{18} m³ requires 1.3×10^{27} discrete grid cells.

And Direct Numerical Simulation (DNS) of the global ocean climate requires 3×10^{10} time steps of one second (1000 years).

So we will be dust long before DNS of global ocean climate is possible:

In the meantime, we must use **subgrid scale parameterizations** to simulate the ocean.

 \rightarrow (See Lesson 4)



Ocean Circulation Models

- ROMS → CROCO <u>https://www.croco-ocean.org/</u>
- NEMO http://www.nemo-ocean.eu/
- MITgcm <u>http://mitgcm.org/</u>
- HYCOM <u>http://hycom.org/</u>
- POP <u>http://www.cesm.ucar.edu/models/cesm1.0/pop2/</u>
- OFES http://www.jamstec.go.jp/esc/ofes/eng/
- MOM http://www.gfdl.noaa.gov/ocean-model
- POM http://www.ccpo.odu.edu/POMWEB/
- etc

Ocean Circulation Models

- Mechanistic studies of ocean and climate processes:
 - Process studies using fine resolution (≤ 1 km) simulations (MITgcm, SUNTANS, CROCO-NH)
 - Mechanisms for coastal and shelf processes (≤ 10 km) (ROMS/CROCO, MARS3D, HYCOM)
 - Mechanisms for climate variability (basin to global) (MOM6, NEMO)
- Operational predictions and state estimation
 - Coastal forecasting India INCOIS
 - Coastal forecasting USA NCEP
 - Ocean state estimation ECCO
- Projections for future climate change
 - IPCC-class simulations with anthropogenic forcing (e.g. CMIP)

Ocean Modelling example:

IPCC global run

Typical length = 100 - 1000 years



0.1 deg

Ocean Modelling example:

Mesoscaleresolving basinscale simulation

Typical length = 10 - 100 years

North-Atlantic coupled simulations:

- oceanic model (6 km)

- atmospheric model (18 km)











Basin-scale configuration **GIGATL**

Submesoscale-permitting basin-scale simulation

Typical length = 1 year

Horizontal resolutions up to **0.75 – 1 km** with 100 topography following levels (refined at the bottom) = 10500 x 14000 x 100 points

including hourly surface forcings and tides.

Run on 10000 processors - 40M cpu-hours - About 4PB of data



Global configuration LLC4320



[[]Torres et al, 19]

Estimate of horizontal grid spacings of the IPCC ocean models

- The gray dots denote the finest grid spacings reported by the IPCC reports by year of publication, except the latest one from the ECCO MITgcm LLC4320 simulation.
- The black line denotes the estimate predicted by Moore's Law, while the shaded regions denote the grid spacing intervals resolving 50% and 90% of surface mixed-layer eddies globally based on the observations and bottom mixed-layer eddies based on simulations.



Ocean Modelling examples:

 $\Delta x = 6 \rightarrow 0.15 \text{ km}$





Submesoscale dynamics at the surface

ROMS ($\Delta x = 500 \text{ m}$)

Submesoscale dynamics at the surface Model Vs Satellite observations

Surface relative vorticity $\frac{\partial v}{\partial x}$

 ∂u

Submesoscale dynamics at the surface Model Vs Satellite observations

Satellite Obs.

Submesoscale dynamics at the surface Model Vs Satellite observations



LATMIX 2012 Campaign (Scalable Lateral Mixing and Coherent Turbulence)

Model

Submesoscale dynamics in the Abyss!



[[]Vic et al., 2018]

Dispersion of larvae by abyssal turbulence



Model vs. Observations

\rightarrow The ocean is a chaotic system



Model vs. Observations

→ The ocean is a chaotic system



24 equi-valid model reproductions of AMOC time series @ 26.5°N.

Model vs. Observations

→ The ocean is a chaotic system





Simulations using ROMS: Regional is a relative concept Animation from J. Molemaker



Normalized relative vorticity, or $Ro = \zeta / f$



Nested domain with open boundaries with dx = 2.5 km

 ζ / f



 ζ / f

dx = 2.5 km



dx = 700 m



dx = 200 m



dx = 50m, **non-hydrostatic**



dx = 50m, **non-hydrostatic**

Surface layer vertical velocity





dx = 15 m (NH)

Surface layer vertical velocity



Activity 1 – Run an idealized ocean basin

Activity 1 – Run an idealized ocean basin SSH



Activity 1 – Run an idealized ocean basin

- Jobcomp (compilation)
- cppdefs.h (Numerical/physical options)
- param.h (gris size/ parallelisation)
- <u>croco.in</u> (choice of variables, parameter values, etc.)

Activity 1 – Run an idealized ocean basin

https://github.com/quentinjamet/Tuto

1) Preparing and compiling the model

For that use the the jobcomp bash file ./jobcomp

- 1. Set library path
- 2. Automatic selection of option accordingly the platform used
- 3. Use of makefile
 - C-preprocessing step : .F → .f using the CPP keys definitions (in cppdefs.h file, customization of the code)
 - Compilation step : $.f \rightarrow .o$ (object) using Fortran compiler
 - Linking step : link all the .o file and the librairy (Netcdf, MPI, AGRIF)
 - --> produce the executable croco

1) Preparing and compiling the model

Edit the param.h and cppdefs.h file to set-up the model

param.h defines the size of the arrays in ROMS:



- Define CPP keys used by the C-preprocessor when compiling the model.
- Reduce the code to its minimal size: fast compilation.
- Avoid FORTRAN logical statements: efficient coding.

1) Preparing and compiling the model

View cppdef.h file

BASIC OPTIONS MORE ADVANCED OPTIONS */ */ /* /* Configuration Name */ # define SOLVE3D # define BENGUELA # define UV COR /* Parallelization */ # define UV ADV # undef OPENMP # ifdef TIDFS # define SSH TIDES # undef MPI # define UV TIDES Embedding */ /* # define TIDERAMP # undef AGRIF # endif Open Boundary Conditions */ /* /* # define CURVGRID # undef TIDES # define SPHERICAL # define OBC EAST # define MASKING # undef OBC WEST /* # define OBC NORTH # define AVERAGES # define OBC SOUTH # define AVERAGES K # define DIAGNOSTICS TS */ Embedding conditions */ # ifdef AGRIF /* # undef AGRIF_OBC_EAST /* # define AGRIF_OBC_WEST /* # define AGRIF OBC NORTH /* * # define AGRIF OBC SOUTH /* # endif /* /* Applications */ /* # undef BIOLOGY /* /* # undef FLOATS # undef STATIONS # undef PASSIVE TRACER # undef SEDIMENTS # undef BBL

Model dynamics */ Grid configuration */ Input/Output & Diagnostics */ # define DIAGNOSTICS UV Equation of State */ ... Surface Forcing */ ... Lateral Forcing */ ... Input/Output & Diagnostics */ ... Bottom Forcing */ ... Point Sources - Rivers */ ... Lateral Mixing */ ... Vertical Mixing */ ... Open Boundary Conditions */ ...

Embedding conditions */ ...

2) Running the model

The namelist roms.in

roms.in provides the run time parameters for ROMS:

title: Southern Benguela time stepping: NTIMES dt[sec] NDTFAST NINFO 5400 60 480 1 S-coord: THETA S, THETA B, Hc (m) 6.0d0 0.0d0 10.0d0 grid: filename Warning ! These ROMS FILES/roms grd.nc should be identical to forcing: filename the ones in ROMS FILES/roms frc.nc bulk forcing: filename romstools_param.m ROMS FILES/roms blk.nc climatology: filename ROMS_FILES/roms_clm.nc boundary: filename ROMS FILES/roms bry.nc initial NRRFC filename 1 ROMS FILES/roms ini.nc NRST, NRPFRST / filename restart: 480 -1 ROMS FILES/roms rst.nc

history: LDEFHIS, NWRT, NRPFHIS / filename T 480 0 ROMS FILES/roms his.nc averages: NTSAVG, NAVG, NRPFAVG / filename 48 0 1 ROMS FILES/roms avg.nc primary history fields: zeta UBAR VBAR U V wrtT(1:NT) T F F F F 10*F auxiliary history fields: rho Omega W Akv Akt Aks HBL Bostr FFFFFFFF primary averages: zeta UBAR VBAR U V wrtT(1:NT) T T T T T 10*T auxiliary averages: rho Omega W Akv Akt Aks HBL Bostr FTTFTFTT rho0: 1025.d0 lateral_visc: VISC2, VISC4 [m^2/sec for all] 0 0 tracer diff2: TNU2(1:NT) [m²/sec for all] 10*0.d0 bottom drag: RDRG [m/s], RDRG2, Zob [m], Cdb min, Cdb max 0.0d-04 0.d-3 1.d-2 1.d-4 1.d-1 gamma2: 1.d0 X SPONGE [m], V SPONGE [m²/sec] sponge: 800 100.e3

nudg_cof: TauT_in, TauT_out, TauM_in, TauM_out [days for all] 1. 360. 10. 360.

Activity 1 - Run an idealized ocean basin



Activity

Homework

- For next time:
 - Read <u>https://www.jgula.fr/ModNum/Stommel48.pdf</u>
 - Read <u>https://www.jgula.fr/ModNum/Munk50.pdf</u>
 - Read

https://elischolar.library.yale.edu/cgi/viewcontent.cgi?article=1817&context=journal_of_marine _research